# Why landslide susceptibility maps should change over time

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# **Objective:**



#### Quantifying the effect of history of landslides in susceptibility modelling

Time-variant landslide susceptibility modelling (Samia et al, 2016, Landslides):

Susceptibility<sub>s,t</sub> =  $f(conditioning attributes_s, previous landslides_{s,t})$ 

Study area, data and inspiration



**Figure 5.** Histograms of geometric and topographic attributes of landslides with (red) and without (blue) follow-up landslides.

Mean of size, roundness, TWI, absolute profile curvature,

ANOVA

relative slope position and vertical distance to channel network between landslides with and with out follow-up landslides are significantly different

Figure 1. Multi-temporal landslide inventory of Collazzone study area in central of Italy





**Figure 2.** Inspiration of the work: many overlapping landslides

Figure 3. Spatial association between landslides

Set of variables	Variables available for logistic regression	AUC calibration	AUC validation
Geometry	2	$0.60 \pm 0.01$	$0.55 \pm 0.03$
Topography	12	$0.57 \pm 0.03$	$0.56 \pm 0.05$
Geometry + topography	14	0.64 ± 0.02	0.58 ± 0.04

**Table 1.** Logistic regression models implemented to predict occurrence of follow-up landslides

## Conclusion



### Results

**Figure 6.** Example: Time-variant landslide susceptibility modelling reflected as clustering of landslides after a first landslide happens within a time scale of about 10 years.

Landslides clearly depend on the history of landsliding, therefore dynamic landslide susceptibility maps are necessary

Elapsed time since first landslide (t) **Figure 4.** Temporal behaviour of landslide path dependency with an exponential decay

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Holomonia fraction p(t) =  $\rho_a \exp^{(b*t)} + \rho_0$   $\rho(t) = \rho_a \exp^{(b*t)} + \rho_0$  $\rho(t) = \rho_a \exp^{(b*t)} + \rho_0$ 

Landslide susceptibility temporarily increases by a factor of 15 following a landslide, and then decreases gradually over time